Modeling Fourier Amplitude Spectra of Earthquake Ground Motions

AECOM

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With thanks to my PhD advisor Norman Abrahamson Jeff Bayless

near-field effects, attenuation, simulations, site response, etc.
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Biography

PhD: University of California, Davis MS, BS: University of California, Los Angeles

Mentors Paul Somerville, Norman Abrahamson, Christine Goulet

Experience in practice is mostly with ground shaking seismic hazard

Research interests include all types of ground motion studies:

Introduction: Response spectra vs. Fourier spectra

Ground Motion Model for FAS

Inter-frequency correlation of the FAS

Response Spectrum

- Is a measure of peak responses of a series of SDOF oscillators different fundamental periods.
- Ubiquitous intensity measure (IM) in earthquake engineering.
- **Because of the SDOF structure, does not directly represent the ground motion itself.**
- **Because it is the peak of the response in time, the response spectrum is not a linear operation.**

https://commons.wikimedia.org/wiki/File:SR-SDoF.gif

Fourier amplitude and phase spectra

- The Fourier Transform is a mathematical model which transforms signals between two domains – e.g. time and frequency.
- Decomposes the signal into sine and cosine components, describing the frequencies present.
- A direct representation of the earthquake ground motion.
- **The Fourier Transform is a linear operation.**

The response spectrum at short periods primarily reflects the strength of the underlying Fourier spectral amplitudes over a broad range of frequencies.

SDOF response (linear DE)

 $u(t) = h(t) \otimes p(t)$

SDOF transfer function FAS of the ground motion (GM)

The response spectrum at short periods primarily reflects the strength of the underlying Fourier spectral amplitudes over a broad range of frequencies.

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Advantages of using FAS in ground motion studies:

The scaling of FAS is easier to constrain using seismological theory, and numerous seismological models of the FAS are available.

(e.g., Brune,1970; Boore, 2004)

Linear site response remains linear at all frequencies and does not depend on the spectral content of the input motion, as is the case for response spectra. *(Stafford et al., 2017)*

In non-ergodic models, which require ample recorded data, effects from small magnitude events are more easily transferred to larger magnitudes.

(Lavrentiadis et al., 2023)

For calibrating finite-fault simulations, Fourier spectra are more closely related to the physics in the simulations.

Covered in part 2 of this talk.

Bayless and Abrahamson (BA18):

empirical ground-motion model (GMM) for shallow crustal earthquakes in California based on the Next Generation Attenuation-West2 database.

For the EAS component of the FAS:

$$
EAS(f) = \sqrt{\frac{1}{2} [FAS_{HC1}(f)^2 + FAS_{HC2}(f)^2]}
$$

The EAS is smoothed in log-space.

 ln EAS = ln EAS_{med} + $\epsilon \sigma$

 \ln EAS_{med} = $f_{Magnitude} + f_{Path} + f_{Site} + f_{Basin} + f_{Depth} + f_{SOF}$

$$
\sigma = \sqrt{\tau^2 + \phi_{S2S}^2 + \phi_{SS}^2}
$$

- Non-linear mixed effects regression independently at each frequency
- § Performed in a series of regression steps
	- **prevents trade-off of correlated model coefficients and to constrain different components of the model using the data** relevant to each
- Emphasis is placed on model scaling and extrapolation outside ranges well constrained by data
- Uses finite-fault simulations to constrain near-fault saturation

$$
f_M = c_1 + c_2(\mathbf{M} - 6) + \frac{(c_2 - c_3)}{c_n} \ln(1 + e^{c_n(c_M - \mathbf{M})})
$$

 $c₂$ is the frequency-independent linear M-scaling slope for frequencies well above the theoretical corner frequency.

The $(c_2 - c_3)/c_n$ term captures both the linear scaling of the FAS below the theoretical corner frequency (coefficient c_3) and the nonlinear transition to that scaling.

The coefficient c_n controls the width of the magnitude range over which the transition between low- and high-frequency linear scaling occurs;

The coefficient c_M is the magnitude at the midpoint of this transition.

$$
f_P = c_4 \ln(R_{\text{rup}} + c_5 \cosh(c_6 \max(\mathbf{M} - c_{hm}, 0))) + (-0.5 - c_4) \ln(\hat{R}) + c_7 R_{\text{rup}}
$$

c4 term: near-source geometric spreading, which is magnitude and frequency dependent.

c5 term: magnitude and frequency dependence on the geometric spreading

- This additive distance is designed to capture the near-source amplitude saturation effects of the finite-fault rupture dimension.
- largest additive distance at high frequencies.

c7 term: crustal anelastic attenuation (CA/Nevada)

 $f_S = f_{SL} + f_{NL} + f_{Z1}$

$$
f_{\rm SL} = c_8 \ln \left(\frac{\min(V_{S30}, 1000)}{1000} \right)
$$

$$
f_{\rm NL} = f_2 \ln \left(\frac{I_R + f_3}{f_3} \right),
$$

$$
f_{Z1} = c_{11} \ln \left(\frac{\min(Z_1, 2.0) + 0.01}{Z_{1\text{Ref}} + 0.01} \right)
$$

 f_{SL} and f_{Z1} are determined empirically from CA/Nevada

 f_{NL} is constrained using a purely analytical model rather than obtaining it from the data (Hashash et al., 2018)

 $\delta_{\text{total}} = \delta B_e + \delta S 2S_s + \delta W S_{\text{es}}$

Near-fault saturation in the model is not as strong as implied by the data; intentionally doesn't allow over-saturation (a peak in distance scaling at some distance greater than zero)

$$
\sigma = \sqrt{\tau^2 + \phi_{S2S}^2 + \phi_{SS}^2}
$$

The standard deviation model is frequency and magnitude-dependent:

At low frequencies, the small-magnitude data have higher between-event standard deviation (τ)

At higher frequencies, not strong magnitude dependence in τ , but ϕ_{SS} and ϕ_{S2S} are larger for the small-magnitude data

Range of Applicability

- 0.1 100 Hz; kappa-based extrapolation beyond 24 Hz
- Regionalized for CA/Nevada, uses global data to constrain the magnitude scaling and geometric spreading
- § Rupture Distances 0 300 km
- § **M** 3-8
- Vs30 from 180 1500 m/s

Future Updates

- Hanging wall model
- Rupture directivity
- Aftershocks
- **Improved site response models**
- **Wider frequency range**
- Partially non-ergodic for other regions
- Non-ergodic (Lavrentiadis et al., 2023)

Inter-frequency correlation of the FAS

Inter-frequency Correlation

Epsilon (ϵ) is the number of standard deviations difference between the observed GM and the median model prediction (ln units)

Inter-frequency Correlation

Repeat this calculation of ρ for each period (or frequency) pair of interest.

Inter-frequency Correlation

Significance of the Inter-frequency Correlation

As a measure of the width of spectral peaks, ρ_{ϵ} is relevant in dynamic structural response.

- Structures are sensitive to a range of frequencies about the fundamental one, especially for nonlinear response.
- ρ_{ϵ} is related to the width of spectral peaks and troughs (higher correlation – wider)
- § Breadth of spectral peaks influences variability of the response.
- Increased ρ_{ϵ} leads to larger response variability, flatter fragility curves (blue), and in turn, higher risk

• Therefore, ρ_{ϵ} **influences seismic risk and is an** important metric for **calibration and validation** of simulations (Bayless and Abrahamson, 2018)

Inter-frequency Correlation in Simulations

CyberShake NZ: crustal earthquake simulations (Bradley et al., 2020)

Evaluation Procedure

- \blacksquare Calculate residuals from the simulations relative to the FAS model
- Partition the residuals into components: mean bias, between-event, between site, within-site
- Calculate ρ_{ϵ} for each FAS residual component and the total ρ_{ϵ}
- Compare with ρ_{ϵ} from BA18

Inter-frequency Correlation: CyberShake NZ

Inter-frequency Correlation: CyberShake NZ

Summary of Performance

The between-site residuals represents the systematic deviation of the observed amplification at a site from the median amplification predicted by the model.

- Generally, are too high in the simulations.
- Impacted by small variability in CyberShake Vs profiles compared with the profiles for the recorded data.
- **Calibration will involve testing simulations with refined velocity models.**

The between-event residuals physically relate to source effects (e.g. stress drop) which drive ground motions over a broad frequency range, and have relatively broad ρ_{ϵ}

- Below about 0.5 Hz performance is good. Above 0.5 Hz the correlation is low.
- **Calibration involves adjustments to source models**

Inter-frequency Correlation: Calibration

Song et al. (2020)

- investigated the effect of pseudo-dynamic source models on the inter-frequency correlation of ground motions with the 1994 Northridge, California, earthquake.
- Pseudo-dynamic emulates the essential physics of earthquake rupture dynamics by statistically analyzing many dynamic rupture models
- The model in (b) have cross correlation between earthquake source parameters (slip, rupture velocity, and peak slip velocity)

Inter-frequency Correlation: Calibration

Song et al. (2020)

- SCEC Broadband Platform **Simulations**
- The affected on ρ_{ϵ} is in the frequency around 0.5 Hz, with little impact on other frequency ranges
- Little impact on the standard deviation
- More work is needed to understand the relation between physics-based earthquake source models and the inter-frequency correlation of ground motions

Summary

The response spectrum at short periods primarily reflects the strength of the underlying Fourier spectral amplitudes over a broad range of frequencies.

As a result, there are numerous advantages to using FAS in ground-motion studies.

One advantage is for calibration and validation of earthquake ground motion simulations, which can be used in future seismic hazard and risk assessments. More research on this topic is needed.

Thank you

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